2	nonstationary characteristics
3	
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12	
13	Abstract
14	The lowest navigable water level (LNWL) is an important indicator used for navigation
15	design to balance the relationship between navigation safety and economic benefits of a
16	waterway. However, it is a challenge of accurately estimating LNWLs due to the
17	nonstationary characteristics of observed water level data series. In this study, a
18	comprehensive framework was developed for handling this issue. In this framework, inter-
19	annual variabilities in both the mean and variance of water level series were described by
20	decomposing original series and were eliminated by composing new series. Intra-annual

A framework for determining lowest navigable water levels with

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21 variability was determined by detecting indicators describing intra-annual water level

22	distributions. Considerations of inter- and intra-annual variabilities were combined by
23	designing annual water level processes for the past and current environments. Shipping risks
24	during both annual and multi-annual periods were considered in the framework as well. The
25	framework was demonstrated in estimating LNWLs at the Gaodao and Shijiao stations in the
26	North River basin, southern China. The recommended LNWLs at the Gaodao station were
27	22.32 m for 95% guaranteed rate and 21.84 m for 98% guaranteed rate; LNWLs at the Shijiao
28	station were 0.27 m for 95% guaranteed rate and 0.15 m for 98% guaranteed rate. The impact
29	of variance variability on estimations of LNWLs was also evaluated. Results indicated that
30	the recommended LNWLs would have errors of 0.11~0.48 m at the Gaodao station and
31	0.03~0.04 m at the Shijiao station if the variance variability was not considered. The proposed
32	framework was then compared with nonstationary synthetic duration curve (NSDC) method,
33	and results illustrated that the duration curves plotted by NSDC method were unreasonable,
34	leading to inaccurate design values. Overall, the developed framework is more reasonable
35	and suitable for designing LNWLs of waterways where the variabilities of the water levels
36	at different time scales are different or where the historical water level data contain various
37	variations.

Keywords: lowest navigable water level; water level processes; nonstationary hydrological
design; intra-annual variability; inter-annual variability; variance variability

#### 42 **1. Introduction**

Inland waterways play an important role in the global transportation system, providing 43 economic benefits with low-cost and environmentally friendly freight transportation modes 44 (Oztanriseven and Nachtmann, 2017; Willems et al., 2018). However, inland waterways are 45 vulnerable to climate variability (Jonkeren et al., 2011; Christodoulou et al., 2020) and human 46 activities (e.g., dam operation and waterway modification) (Valle and Kaplan, 2019) as 47 48 waterway capacities largely depend on the water levels of rivers (Wang et al., 2020). For example, in the dry period waterways are adversely affected by the resultant lowered water 49 levels (Linde et al., 2017). Since the navigation safety is threatened in such situation, the load 50 amounts must be restricted or shipping must be interrupted to wait for the water levels to rise 51 again (Jonkeren et al., 2014), leading to increases in transport costs and decreases in shipping 52 efficiency. 53

To guarantee efficient inland navigation, concepts related to navigable depth/water level 54 are adopted with which to define safe and effective channel bottom criteria. Among them, the 55 lowest navigable water level (LNWL) has been used as an important navigation standard for 56 inland waterways (Zhao et al., 2018; Yang et al., 2019), especially in China. However, the 57 design of LNWLs requires long-term water level series, which usually show characteristics 58 of nonstationarity under changing environments (Milly et al., 2008). These nonstationary 59 characteristics appear not only in water levels at different timescales (e.g., annual, seasonal, 60 61 dry or wet period, month, and day) but also in intra-annual distributions of water levels (Shiau and Wu, 2007; He et al., 2019; Yao et al., 2020). Under this circumstance, the LNWLs 62

designed in the past or based on stationary assumptions may be unreasonable. An
overestimated LNWL can reduce economic benefits, as the potential of the waterway
capacity is not fully exploited. If the LNWL is underestimated, the risks of ship grounding
and the costs of dredging (Kling et al., 2003) will increase. Therefore, methods for designing
reasonable LNWLs under conditions of nonstationarity are required.

Presently, many approaches for nonstationary hydrological frequency analysis have 68 69 been proposed, and they include two main types: direct estimation approaches using nonstationary probabilistic models, and indirect estimation approaches with reconstructions 70 of original series (Liang et al., 2017; Zhao et al., 2018; Feng et al., 2020). Among them, 71 mixed distribution models (e.g., Singh and Sinclair, 1972; Waylen and Woo, 1982; Yan et al., 72 73 2016), time-varying moments models (e.g., Strupczewski et al., 2001a, 2001b; Strupczewski and Kaczmarek, 2001; Villarini et al., 2009; Cannon, 2009; Vogel et al., 2011; Jiang et al., 74 75 2015; Liu et al., 2014; Ahn and Palmer, 2016; Li et al., 2019), and conditional probability distribution models (e.g., Singh et al., 2005) are typical models in the first type, and these 76 models have been widely used and improved. However, previous studies (e.g., Montanari 77 and Koutsoyiannis, 2014; Milly et al., 2015; Zheng et al., 2018; Jiang et al., 2019) have 78 79 pointed out that these models may have too many parameters to estimate, resulting in estimation errors and uncertainties (especially when input data series are short). Additionally, 80 the return periods and risk estimates for nonstationary conditions are different from those 81 82 corresponding to stationary conditions, and the design values can thus be uncertain for a given exceedance probability (Salas and Obeysekera, 2014; Rosner et al., 2014; Jiang et al., 83

2015; Sarhadi et al., 2016). Although some risk-based nonstationary design strategies such
as "expected waiting time" (Salas and Obeysekera, 2014), "design life level" (Rootzen and
Katz, 2013) and "equivalent reliability" (Hu et al., 2018) have been proposed for estimation,
the reliability of the final design values may need further improvement by considering
uncertainties.

The second type of approaches requires reconstructing the historical series to meet the 89 90 requirement of "stationarity". These approaches assume that a nonstationary time series is composed of nonstationary deterministic parts (Maidment, 1993) and a stationary stochastic 91 part (Yevjevich, 1972; Guttman and Plantico, 1989; Milly et al., 2015; Stojković et al., 2017). 92 Therefore, in these methods, the original series is first decomposed into different parts and 93 then a new same-length stationary series is composed for estimations. For example, Gau et 94 al. (2007) decomposed groundwater level series into three components by the additive model 95 96 performed by STATISTICA (Statsoft Inc., 2003). Wavelet analysis is also used to decompose and reconstruct hydrological series by wavelet transform (Wang et al., 2015). Hu et al. (2015) 97 proposed the concept of expected vibration center (EVC), separated the original series by a 98 novel optimal segmentation technique, and used the EVC to reconstruct the original 99 100 hydrological series. Gado and Nguyen (2016a, 2016b) removed trends in both the mean and standard deviation from the original series to obtain transformed "stationary" time series for 101 subsequent calculations. Liang et al. (2017) proposed a modified reservoir index (MRI(t)) 102 103 and reconstructed annual maximum flow series by multiplying the scalar factor 1/(1 - MRI(t)). Ren et al. (2018) used seasonal-trend decomposition based on loess (STL) (Cleveland et al., 104

105 1990) to decompose the components of monthly streamflow series and then recombined them.
106 Zhao et al. (2018) used several methods to detect abrupt changes, trends and periodicities in
107 observed series, and then used an additive model to compose a new series with a time108 invariant mean.

109 Compared to the first type, the second type of approaches has relatively fewer parameters for estimations. Another theoretical advantage is that identifying deterministic 110 111 components of historical series allows for a better interpretation of the temporal variability present in the series and the physical causes of the nonstationarity. However, this type of 112 approaches mainly concerns the variability in the mean, but lacks deep considerations of 113 abrupt changes, trends and periodicities in the variance and high-order moments. Overall, 114 both types of approaches have not been widely used to analyze nonstationary water levels, 115 let alone to design LNWLs with nonstationary characteristics. 116

117 The nonstationary synthetic duration curve (NSDC) method has been proposed for designing LNWLs (Zhao et al., 2018), which belongs to the second type described above. 118 However, this method has some limitations. First, it only detects and deals with the inter-119 annual variability in annual average water level series. However, water levels at different 120 timescales during the year may respond differently to the changing environment (López-121 Moreno et al., 2013; Palleiro et al., 2014), and therefore variabilities in low water levels can 122 be exaggerated or overlooked during the design, leading to unreasonable LNWL estimations. 123 124 Besides, the navigable guaranteed rate used in the NSDC method only reflects the daily average navigable probability during the multi-year period, lacking the consideration of 125

126 annual shipping risks. Furthermore, some studies (e.g., Tu et al., 2015; Ren et al. 2018; Cui et al., 2020) have evaluated indicators reflecting the characteristics of intra-annual flow 127 regime alterations such as the complete accommodation coefficient ( $C_c$ ), concentration 128 degree  $(C_d)$  and non-uniformity coefficient  $(C_n)$ , and have pointed out the existence of intra-129 annual variability in streamflow. Whereas the variability in distributions of intra-annual water 130 levels was not considered in the NSDC method either, leading to inappropriate duration 131 132 curves and unreasonable estimations of LNWLs. Therefore, how to handle the intra-annual variability of water level and how to combine the considerations of inter- and intra-annual 133 variabilities in the design process require further study. 134

The objective of this paper is, therefore, to build a more comprehensive framework for 135 designing the LNWL. This framework consists of the following improvements to the NSDC 136 method. First, considering the advantages of the second type of approaches listed above, the 137 idea of decomposition and composition is also employed in this study, and the deficiency of 138 only handling mean variations in most previous studies is overcome. Thus, reconstructed 139 series with time-invariant means and variances can be obtained for subsequent design. 140 Second, the variabilities in annual water level process, including variabilities in low water 141 levels and intra-annual water level distributions, are considered together. By considering 142 these variations, the impact of nonstationarity on the estimations of LNWLs can be 143 minimized. Another advantage of this framework is that it considers not only the annual 144 145 shipping risks occurring in the multi-year duration (expressed by the "exceedance probability"), but also the guaranteed navigable days at annual scales (expressed by the 146

"annual navigable guaranteed rate"). The application of the framework is demonstrated in
estimating the LNWLs at Gaodao and Shijiao hydrological stations located in the North River
Basin, southern China. The daily water level data used from 1955 to 2016 at Gaodao station
and from 1952 to 2017 at Shijiao station are used for this study.

# 151 2. The proposed framework for designing LNWLs

The proposed framework of designing LNWLs includes the following parts (see Fig. 1): 152 153 (1) Water levels at multi-temporal scales are detected to identify the inter-annual variabilities in the means and variances, and each series is decomposed into abrupt changes, trends, 154 periodic components, and stochastic remainders; (2) Each water level series is reconstructed 155 under the assumed environment. Then, designed water levels at multiple time scales for a 156 given "exceedance probability" are calculated based on the new series. (3) Annual series of 157  $C_d$ ,  $C_n$ ,  $C_c$ , and GI, which characterize the intra-annual water level variabilities, are calculated 158 and detected. (4) All the change-points are used to split the time period, and the typical year 159 of each sub-period is selected. (5) The designed intra-annual water level hydrograph under 160 each environment is derived by scaling the benchmark water level hydrograph of the typical 161 year. (6) The duration curve is drawn based on the designed intra-annual water level 162 hydrograph, and the LNWL is estimated for a given annual navigable guaranteed rate. 163

164 <Figure 1>

# 165 **2.1 Characterization of intra-annual water level distribution**

166 The concentration degree  $(C_d)$ , non-uniformity coefficient  $(C_n)$ , and complete 167 accommodation coefficient  $(C_c)$  are the indicators commonly used for characterizing the 168 intra-annual flow regime alterations. In this section, these indicators along with the Gini coefficient (GI) are used to describe the characteristics of intra-annual water level 169 distributions. The larger the  $C_d$ ,  $C_n$  and  $C_c$  are, the more uneven the intra-annual water level 170 process is. However, GI is negatively correlated with the non-uniformity. By detecting these 171 indicators simultaneously, the intra-annual variability of water levels can be reliably obtained. 172  $C_d$  is calculated based on the premise that the daily water level is a vector with a unique 173 174 direction and magnitude. The direction represents the corresponding circular date and the magnitude is the value of the water level (Gumbel, 1954; Magilligan and Graber, 1996; 175 Magiligan and Nislow, 2005). Therefore, the calendar date (i) can be transformed into the 176 circular date  $(\theta)$ , such that: 177

$$\theta_i = \frac{2i-1}{2} \times \frac{360^\circ}{n} \tag{1}$$

179 where *n* is the number of days in the year,  $\theta_i$  is the circular date of the *i*<sup>th</sup> day, and *i* ranges 180 from 1 to *n*.  $C_d$  is expressed as follows:

181 
$$C_d = \frac{Z}{\sum_{i=1}^n Z_i}$$
(2)

182 
$$Z = \sqrt{\left(\sum_{i=1}^{n} Z_{i} \cos \theta_{i}\right)^{2} + \left(\sum_{i=1}^{n} Z_{i} \sin \theta_{i}\right)^{2}}$$
(3)

183 where Z is the module of the resultant vector of all daily water levels; and  $Z_i$  is the 184 magnitude of the water level of the  $i^{th}$  day.

185  $C_n$  can measure the dispersion of data points around the mean value and is calculated as:

186  $C_n = \frac{\sigma}{\overline{Z}}$ (4)

187 
$$\overline{Z} = \frac{1}{n} \sum_{i=1}^{n} Z_i$$
(5)

188 
$$\sigma = \sqrt{\frac{1}{n} \sum_{i=1}^{n} (Z_i - \overline{Z})^2}$$
(6)

189 where  $\overline{Z}$  is the annual average water level, and  $\sigma$  is the standard deviation of the daily 190 water level series. The formula of  $C_c$  is shown below:

191 
$$C_{c} = \frac{\sum_{i=1}^{n} \phi(i)(Z_{i} - \overline{Z})}{\sum_{i=1}^{n} Z_{i}}$$
(7)

192 
$$\phi(i) = \begin{cases} 0, Z_i < \overline{Z} \\ 1, Z_i \ge \overline{Z} \end{cases}$$
(8)

*GI* has been widely used to analyze spatial-temporal variabilities in hydrological process
(Jawitz and Mitchell, 2011; Shi et al., 2013; Masaki et al., 2014; Cai et al., 2018). Here, it is
used to characterize the intra-annual water level distribution. The Lorentz curve is drawn by
taking the cumulative percentage of days as the independent variable and the cumulative
percentage of daily water levels as the dependent variable. *GI* is expressed as:

$$GI = \frac{S_B}{S_A + S_B} \tag{9}$$

199 where  $S_A$  and  $S_B$  are the areas of sections A and B (Fig. 2), respectively.

200 <Figure 2>

# 201 2.2 Comprehensive detection for hydrological variability

The deterministic components in the water level series are first detected by using statistical methods. After subtracting the deterministic components of the mean from original series, we square each value in the residual series. Then, the same detection process is applied to the squared residual series (Vinnikov and Robock, 2002) to detect the deterministic
components of the variance. The detection/decomposition process is presented as follows.

207 An initial judgment about the stationarity/nonstationarity of the series is made by the 208 ranked version of the von Neumann's ratio (*RVN*) test (Bartels, 1982) at the significance level 209 of  $\alpha$  (Machiwal and Jha, 2012). The *RVN* formula is given below:

210 
$$RVN = \frac{\sum_{i=1}^{n-1} (R_i - R_{i+1})^2}{\sum_{i=1}^n (R_i - \overline{R})^2}$$
(10)

211 where,  $\overline{R} = \frac{1}{n} \sum_{i=1}^{n} R_i$ ,  $R_i$  is the rank of the *i*<sup>th</sup> observation, and *n* is the length of the samples.

If the series shows nonstationarity, its trend is detected by using the nonparametric Mann-Kendall test (Mann, 1945; Kendall, 1955) and is described using a linear equation:

214 
$$y_{trend} = a + bt \tag{11}$$

where parameters a and b can be estimated using the least squares fitting method.

After that, the Moving T test (Afifi and Azen, 1972), Brown-Forsythe test (Brown and 216 Forsythe, 1974) and Lee-Heghinian (Lee and Heghinian, 1977) method are used to detect 217 abrupt changes. Since the detected change-points are sometimes rather different due to the 218 methods used (Li et al., 2016; Xie et al., 2018), the results of the three methods need to be 219 considered comprehensively. If two (or all) of the three methods obtain the same change-220 point, this change point is confirmed. If the change-points obtained by the three methods are 221 222 completely different, then a comprehensive judgment should be made based on investigated physical causes. The detected abrupt change is expressed as: 223

224 
$$y_{abrupt} = \begin{cases} 0, t = 1, 2, ..., n_1 \\ \overline{x_{in2}} - \overline{x_{in1}}, t = n_1 + 1, n_2 + 2, ..., n \end{cases}$$
(12)

where  $n_1$  and  $n_2$  are the lengths of sub-series before and after the change-point, respectively,  $\overline{x_{m1}}$  and  $\overline{x_{m2}}$  are the means of the sub-series before and after the change-point, respectively. The significant trend or abrupt change should be subtracted from the original series. When both forms show significance, the *R*-squared value is introduced to determine the major form of variability, which is given as follows:

230 
$$R^{2} = 1 - \frac{\sum_{i=1}^{n} (Q_{obs,i} - Q_{sim,i})^{2}}{\sum_{i=1}^{n} (Q_{obs,i} - \overline{Q}_{obs})^{2}}$$
(13)

where  $Q_{obs,i}$  is the observed hydrological series with the length n,  $\overline{Q}_{obs}$  is the mean of the observed hydrological series, and  $Q_{sim,i}$  is the sum of the assumed deterministic component and the stochastic component. The trend or the abrupt change with the larger  $R^2$  should be removed from the series in priority.

Then, the *RVN* test is applied to the residual series. When the residual is nonstationary, then the test is continued to check if the series contains another trend or abrupt change. After all the trends and abrupt changes are subtracted, the Fourier series (Bras and Rodriguez-Iturbe, 1993) is used to detect the periodic component (Machiwal and Jha, 2012) if the remainder still shows nonstationarity.

240 
$$y_{periodicity} = A_0 + \sum_{k=1}^{l} [A_k \sin(\frac{2\pi kt}{T}) + B_k \cos(\frac{2\pi kt}{T})]$$
(14)

where  $A_0$  is the population mean, *l* is the total number of harmonics (l=T/2 for even *T* and l=(T+1)/2 for odd *T*), *T* is the base period or period of the function,  $A_k$  and  $B_k$  are the sine 243 and cosine Fourier coefficients, respectively.

#### 244 2.3 Composing stationary water level series

The deterministic components  $y_t$  (including  $y_{trend}$ ,  $y_{abrupt}$  and  $y_{periodicity}$ ) of each water level time series are identified by the methods described above. Then, the stochastic component  $s_t$  can be obtained as:

248 
$$s_{t} = x_{t} - y_{t} = \frac{x_{t} - a_{t}}{b_{t}}$$
(15)

249 where  $x_t$  is the original series,  $a_t$  and  $b_t$  are the deterministic components of the mean 250 and standard deviation, respectively. 251 The composed water level series  $x_{t,t_0}$  is a combination of the stochastic component  $s_t$ 

and deterministic component  $y_{t_0}$  under the environment of year  $t_0$ . The composition formula is shown as follows:

254 
$$x_{t,t_0} = a_{t_0} + b_{t_0} s_t = a_{t_0} + \frac{b_{t_0}}{b_t} (x_t - a_t)$$
(16)

where  $x_{t,t_0}$  is the composed water level series under the environment of year  $t_0$ , constants a<sub>t\_0</sub> and b<sub>t\_0</sub> are deterministic components of the mean and standard deviation under the environment of year  $t_0$ , respectively. Based on the reconstructed series, Pearson-III frequency curves are used to estimate annual 3-, 10-, 15-, 30-, 60-, 90-, 120-, 150-, 180-, 240-, 300-day minimum water levels and annual average water level for a given exceedance probability.

#### 261 2.4 Design of intra-annual water level hydrograph and LNWLs

First, the change-points of both inter- and intra-annual variabilities are used to split the

263 whole time period. The changing environment is therefore divided into several relatively stable stages, because the water level condition in each sub-period is relatively stable. It 264 should be noted that there is no need to split the time period if no abrupt change is identified. 265 Second, typical years in different sub-periods/environments are selected. Abnormal 266 hydrological years should be excluded before this selection, because the water level process 267 of the selected typical year should fully reflect the water level condition of the corresponding 268 269 environment. During this selection, the worst water level situation for shipping is considered to ensure safe and continuous navigations. Specifically, the following principles are followed 270 to obtain reasonable typical years and benchmark water level hydrographs: (1) The annual 271 3-, 10-, 15-, 30-, 60-, 90-, 120-, 150-, 180-, 240-, 300-day minimum water level values of the 272 typical year should be close to the corresponding designed water levels; (2) the duration for 273 which the water levels are below the multi-year average water level should be long; and (3) 274 275 when the chosen years follow the first two principles, the year with a relatively large  $C_n$  is preferentially chosen. 276

The third step is to design the intra-annual water level process for assumed environment by scaling the benchmark water level hydrograph of the typical year. Different amplifiers are calculated by Eq. (17). This equation is sourced from a method for designing flood processes and ensures the water level values at different time scales in the designed water level process have the same exceedance probability.

$$K_{j} = \begin{cases} \frac{Z_{d,j}}{Z_{t,j}}, & j = 1\\ \frac{Z_{d,j} - Z_{d,j-1}}{Z_{t,j} - Z_{t,j-1}}, & j \ge 2 \end{cases}$$

$$14$$
(17)

where  $K_j$  is the amplifier for the  $j^{\text{th}}$  time scale,  $Z_{d,j}$  and  $Z_{d,j-1}$  are the designed water levels for the  $j^{\text{th}}$  and  $(j-1)^{\text{th}}$  time scales, respectively,  $Z_{t,j}$  and  $Z_{t,j-1}$  are the actual water levels of the typical year for the  $j^{\text{th}}$  and  $(j-1)^{\text{th}}$  time scales, respectively.

286 <Figure 3>

In the fourth step, daily average water levels in the designed water level hydrograph are sorted from large to small, and then they are subdivided according to the percentages of time during which specific water levels are equaled or exceeded. After that, the relationship between the percentage of time (i.e., guaranteed rate) and water level is drawn as the water level duration curve. Finally, the designed LNWL can be obtained from the duration curve according to the annual navigable guaranteed rate.

#### 293 **3. Study area and data**

The North River is the second largest tributary river in the Pearl River system in South 294 China (Figure 4). It is a main waterway connecting the southern and northern parts of 295 Guangdong Province and has a great significance in transportation and trading (Luo et al., 296 2008; Li et al., 2019). However, various navigation standards for the river, including the 297 LNWL, need to be redesigned. One reason is that the waterway is planned to upgrade from 298 grade VI to grade III to satisfy the increasing demand of waterway capacity of the North 299 River. Another reason is that human activities have changed the streamflow characteristics in 300 the North River (Du et al., 2020). For example, the construction and operation of the Feilaixia 301 302 hydroelectric project (see Figure 4) have caused the upstream area to be affected by the backwater since March 1999 (Wu et al., 2014), and the navigation condition in the upper 303

304 river has improved significantly. The water level in the downstream area also changed due to dam closure and reservoir regulations. Additionally, Lu et al. (2007) have measured annual 305 channel cross-sections at Sanshui hydrological station in the lower North River, and pointed 306 out that extensive sand mining in the lower North River has caused rapid channel incision 307 and water level reduction. Dai et al. (2008) analyzed temporal variations in sediment load at 308 Shijiao station in the North River and found that the sediment flux of the station has reduced. 309 310 The reduction of sediment load is due to effective soil and water conservation projects, dam constructions and regular channel dredging in upstream areas, and such human activities are 311 likely to contributed to bed downcutting to some extent as well (Lu et al., 2007). Overall, the 312 LNWL used in the past cannot satisfy the current conditions anymore. 313

#### 314 **<Figure 4>**

In this study, the observed daily average water level series at Gaodao station from 1955 315 to 2016 and at Shijiao station from 1952 to 2017 are used to illustrate the proposed framework. 316 The two stations comply with the highest hydrological monitoring standards in China, and 317 are located in the middle and low reaches of the North River, respectively. Although the 318 distance between the two stations is not far (xx km), the water level variations of the two 319 stations are quite different. It should be noted that the disruptions of navigation in the area 320 are caused by low water levels. The navigation disruptions caused by ice formation do not 321 exist in this study area, due to its subtropical warm climate conditions. 322

#### 323 **4. Results**

#### **4.1 Inter- and intra-annual variability in the water level**

Hydrological year in this region starts from April 1<sup>st</sup> and lasts until March 31<sup>st</sup> of the following year. The annual minimum water levels at multiple scales and different indicators are calculated under the hydrological year. Then, the annual 3-, 10-, 15-, 30-, 60-, 90-, 120-, 150-, 180-, 240-, 300-day minimum water level time series and annual average water level time series are detected, and the results are shown in Table 1 (More details can be found in Appendix A).

331 <Table 1>

332 **<Figure 5>** 

333 **<Figure 6>** 

On the whole, the water level variations at different time scales are roughly similar, but 334 some change-points and the degrees of variations are different. In terms of the variability in 335 the first-order moments, all the water level series at Gaodao station have abrupt upward 336 changes, and the change-points in extremely low (the 3-, 10- and 15-day minimum) water 337 level series occur in 1999. While all the water level series at Shijiao station have abrupt 338 downward changes, and the change-points in the low water level series occur in 2006. With 339 regard to the variability in the second-order moments, only the 3-, 10-, and 15-day minimum 340 water level series at Gaodao station contain change-points. The change-point of 3-day 341 342 minimum water level series occurs in 1999, and the change-points of 10- and 15-day minimum water level series both occur in 2002. At Shijiao station, almost all the water level 343

series have two change-points in their second-order moments, occurring in 2003 (abrupt
upward changes) and 2007 (abrupt downward changes). Among the series, the 3-, 15-, 120day minimum water level series and annual average water level series are taken as examples
(Figs. 5-6) to intuitively display the inter-annual water-level variabilities at Gaodao and
Shijiao stations.

For Gaodao station, the construction and operation of the Feilaixia water conservancy 349 350 project on the main stream of the North River are regarded as the main causes of variabilities. The river closure work was started in 1998 and the whole Feilaixia project was completed in 351 1999, and these two time-nodes are consistent with most of the change-points above. In 352 addition, the abrupt upward changes of various water level series are consistent with the fact 353 that the river water level at Gaodao station sharply increased due to the retaining effect of the 354 dam. Moreover, the relationship between low water levels and low flows has become 355 complicated and irregular, which corresponds to the abrupt upward changes in the variances 356 of low water level series. In terms of Shijiao station, the water level was affected by 357 regulations of upstream Feilaixia reservoir, high flows in downstream confluence stream, and 358 local excessive sand mining. These factors made the water-level variability complicated. 359 Caused by excessive sand mining, the annual average water level and low water levels of 360 Shijiao station continually dropped significantly from 2003 to 2010, which was similar to the 361 detection results. 362

To detect the variability in intra-annual water level distribution, annual  $C_d$ ,  $C_n$ ,  $C_c$  and *GI* are calculated and their series are used for detection. The detection results of intra-annual 365 variability are shown in Table 2.

366 <**Table 2**>

The intra-annual water level variabilities at Gaodao and Shijiao stations differ greatly 367 due to their different positions in the North River basin. All the indicators of Gaodao station 368 have abrupt changes  $(C_d \downarrow, C_n \downarrow, C_c \downarrow \text{ and } GI^{\uparrow})$  in 1998. It means that the intra-annual water 369 level process became more homogeneous (Fig. 7(a)) after 1998, which is in accordance with 370 371 the retaining effect of dam after 1998. The series of  $C_n(\uparrow)$ ,  $C_c(\uparrow)$  and  $GI(\downarrow)$  at Shijiao station have abrupt changes in 2007, and the change-point in the series of  $C_d(\uparrow)$  is 2006. Accordingly, 372 the intra-annual water level hydrograph at Shijiao station is more uneven than that in the past 373 (Fig. 7(b)). According to the investigation, the water level in the flood season could be lifted 374 more obviously by backwater flooding because of the riverbed downcutting at Shijiao station, 375 resulting in sharper and multi-peak water level processes. 376

377 <Figure 7>

# 378 **4.2 Design of the LNWLs**

There are three types of change-points related to the intra-annual variability, to the interannual variability in the mean and to the inter-annual variability in the variance, respectively. After excluding recurring change-points, all the change-points that appear at Gaodao station occur in 1998, 1999, and 2002. However, since 1998 and 1999 are rather close, they should be merged into one point when splitting the time series. Here we choose 1998 when the Feilaixia project started as the merged change-point. As a consequence, three sub-periods, i.e., 1955 to 1998, 1999 to 2002 and 2003 to 2016 are obtained. The first sub-period

386	represents the past environment in which there is no human-induced water level variation;
387	the transition period from 1999 to 2002 contains a variety of water level variations due to
388	external impacts; and the post-impact period from 2003 to 2016 stands for the current
389	environment where all kinds of water level variations have already happened. At Shijiao
390	station, the change-points occur in 2003, 2005, 2006 and 2007. Similarly to the sub-periods
391	created for Gaodao station, the entire time period at Shijiao station is also divided into three
392	parts: the pre-impact period from 1953 to 2003, the transition period from 2004 to 2007 and
393	the post-impact period from 2008 to 2017.
394	Then, water level series are reconstructed for the pre- and post-impact periods, and the
395	water levels $Z_{d,j}$ are estimated based on reconstructed series by using Pearson-III frequency
396	curves. The estimated parameters are shown in Table 3, and the fitting efficiency of each
397	water level series is higher than 95%.

398 <**Table 3**>

399 <Table 4>

Since the North River is a grade-III waterway, the return period for the disruption of shipping should be 4 years or 5 years according to the national navigation standard (GB50139-2014) of inland waterways in China (Table 4). Here, the return period is assumed to be 5 years (i.e., exceedance probability is 80%). Therefore, the 3-, 10-, 15-, 30-, 60-, 90-, 120-, 150-, 180-, 240-, 300-day minimum water levels and annual average water levels designed at the exceedance probability of 80% (Table 5) are used for the subsequent design. **<Table 5>** 

407	Next, some years are selected as candidate for typical years (Appendix B) according to
408	the principles mentioned in Section 2.4. Considering them comprehensively, 1983 and 2004
409	are selected as typical years of Gaodao station, and their water level hydrographs represent
410	the benchmark water level processes of pre- and post-impact periods, respectively. The years
411	1987 and 2013 are selected as typical years for the pre- and post-impact periods at Shijiao
412	station, respectively. Then the designed intra-annual water level processes for both periods
413	(Figs. 8(a) and (b)) are obtained by scaling the corresponding benchmark hydrographs (Table
414	5). Based on the designed water level processes, water level duration curves of the pre- and
415	post-impact periods (Figs. 8(c) and (d)) are drawn. Finally, the LNWLs at the annual
416	navigable guaranteed rates of 95% and 98% are obtained from the duration curves. The
417	recommended LNWLs at Gaodao station for the current environment are 22.32 m for 95%
418	guaranteed rate and 21.84 m for 98% guaranteed rate; LNWLs at Shijiao station for the
419	current environment are 0.27 m for 95% guaranteed rate and 0.15 m for 98% guaranteed rate.
420	The designed water level hydrograph of Gaodao station for the post-impact period (i.e.,
421	current environment) is more uniform than the hydrograph for the pre-impact period (i.e.,
422	past environment). Compared to the past, the designed water level hydrograph of Shijiao
423	station for the current environment contains more obvious seasonal changes. Both
424	conclusions are consistent with the detected results of intra-annual variabilities. It can also
425	be seen in Fig. 8(c) that the tail of the duration curve of the post-impact period has a larger
426	slope than that of the pre-impact period, implying that the intra-annual fluctuation of low
427	water levels in the current environment is more intense than that in the past environment at

428	Gaodao station. However, at Shijiao station, the intra-annual fluctuation of low water levels
429	in the current environment is similar to that in the past, because the tails of the two duration
430	curves are approximately parallel (Fig. 8(d)). Besides, the differences in the LNWL for the
431	same guaranteed rate between the past and current environments are 2.48 m (for 95%
432	guaranteed rate) and 2.05 m (for 98% guaranteed rate) at Gaodao station, -4.14 m (for 95%
433	guaranteed rate) and -4.21 m (for 98% guaranteed rate) at Shijiao station, respectively. These
434	differences indicate the necessity of considering nonstationarity when designing the LNWL.
435	<figure 8=""></figure>
436	5. Discussion
437	5.1 Impact of variance variability on estimating LNWLs
438	The LNWLs estimated by the proposed framework without considering the variability
439	in variance are used for comparison, and the differences between the LNWLs estimated under
440	different considerations are used to evaluate the impact of variance variability on LNWLs.

441 The LNWLs estimated under different considerations are shown in Table 6.

For Gaodao station, differences in the LNWLs caused by considering and not considering the variance variability are 0.03 m (for 95% guaranteed rate) and 0.02 m (for 98% guaranteed rate) in the past environment. In the current environment, the differences are 0.11 m when the guaranteed rate is 95% and 0.48 m when the guaranteed rate is 98%, accounting for approximately 6%~30% of the ship draft of standard ships in the North River. If the variance variability is not considered, the designed value will be overestimated. A higher LNWL means a smaller capacity of the waterway. In that case, the LNWL supposed to have 95% guaranteed rate actually has 93% guaranteed rate due to the impact of variance
variability, and the LNWL supposed to have 98% guaranteed rate actually has 95%
guaranteed rate. Therefore, the guaranteed rate for navigation at Gaodao station will be
artificially reduced, the waterway will not be fully exploited and the shipping benefits will
be reduced.

For Shijiao station, differences between the LNWLs estimated with and without 454 455 consideration of the variability in variance are -0.17 m (for 95% guaranteed rate) and -0.19 m (for 98% guaranteed rate) in the past environment. Differences in the LNWLs of the 456 current environment are 0.04 m when the guaranteed rate is 95% and 0.03 m when the 457 guaranteed rate is 98%, accounting for 14.8% and 20% of the corresponding recommended 458 LNWLs, respectively. At Shijiao station, if the variance variability is not considered, the 459 designed value will be underestimated. To prevent ships from bottoming out, there should be 460 a certain under keel clearance (UKC) in the channel. Therefore, there will have a greater risk 461 of ship grounding. The underestimation seems not significant in the current environment, 462 however, since the water level itself at Shijiao station is relatively low, a few centimeters of 463 estimation error equals to 15%~20% of the design result. In that case, the increased shipping 464 risk should not be ignored. 465

466 **<Table 6>** 

467 **<Figure 9>** 

As shown in Fig. 9, there are obvious differences between the reconstructed low water level series (the 3-, 10-, and 15-day minimum water level series) with and without

470	consideration of the variance variability. These differences can affect the designed water
471	levels $\overline{Z_{d,j}}$ and the amplifiers for the design of water level processes, so as to influence the
472	estimation of the LNWL. At Gaodao station, the designed 3-, 10-, and 15-day minimum water
473	levels in the current environment are 21.52, 21.79 and 21.82 m (Table 5). When the variance
474	variability is not considered, the corresponding designed values change to 21.89, 22.25 and
475	22.32 m. At Shijiao station, the designed 3-, 10-, and 15-day minimum water levels in the
476	current environment are 0.13, 0.24 and 0.32 m (Table 5). The corresponding designed values
477	are 0.17, 0.28 and 0.34 m when the variance variability is not considered. Obviously, at
478	Shijiao station, the differences of the designed water levels with different considerations are
479	relatively small, which explains why the impact of variance variability on LNWLs at Shijiao
480	station is not significant.

For further analysis, the tails (guaranteed rate is larger than 80%) of water level duration 481 curves under different considerations are compared using Tukey-boxplots (Sheskin, 2011). 482 The minimum values and lower quartiles of low water levels from two duration curves have 483 clear differences in Figs. 10(b) and (c), and the differences between the LNWLs in both cases 484 are also significant. Whereas in Figs. 10(a) and (d), the minima (including the outliers 485 486 denoted in red points) and lower quartiles under the two considerations are relatively close, and in the past environment of Gaodao station and the current environment of Shijiao station, 487 the differences between the LNWLs are not very significant either. Therefore, the influence 488 489 of variance variability on LNWLs basically depends on its influence on low water levels when reconstructing the series. 490

491 <Figure 10>

For Gaodao station, the pre-impact period is 30 years longer than post-impact period. 492 When the variance variability is considered in the reconstruction of the series for the past 493 environment, a relatively small number of low water levels are corrected. When the series 494 before the change-point is modified to the current environment, a relatively large number of 495 water level values become lower due to the variance variability, and therefore the low water 496 497 level part of the whole series becomes lower. However, for Shijiao station, although the preimpact period is 41 years longer than the post-impact period, when reconstructing the series 498 for the past environment, many low water level values in the transition and post-impact 499 periods which far away from the mean are corrected, resulting in a big impact on the low 500 water level part of the whole series. Conversely, in the current environment, whether the 501 variance variability is considered or not has less influence on the low water level part. 502

503

# **5.2** Comparison with the NSDC method

In this section, the proposed design framework for LNWLs considering both the inter-504 and intra-annual variabilities in water levels is compared with the NSDC method mentioned 505 above. In the NSDC method, the annual average water level is used to represent the overall 506 water level condition each year. Thus, only the annual average water level series is 507 reconstructed for the past and current environments. Because the form and extent of the 508 variability on each day of the year are regarded as the same, stationary daily water level 509 510 series for each year are obtained by multiplying the observed daily water level series with the corresponding annual scaling ratio. Finally, all the daily water level values are combined 511

to plot a multi-year synthetic duration curve (Figs.11 (a) and (b)), and the LNWL isestimated according to the multi-year guaranteed rate.

514 <Figure 11>

515 **<Table 7>** 

As shown in Table 7, LNWLs of the current environment at Gaodao station estimated 516 by the NSDC method are 22.25 m for the guaranteed rate of 95% and 22.08 m for the 517 518 guaranteed rate of 98%. The differences in the LNWLs between two methods at Gaodao station are -0.07 m (95% guaranteed rate) and 0.24 m (98% guaranteed rate). LNWL of 519 Shijiao station estimated by the NSDC method under the current environment is 0.59 m for 520 the guaranteed rate of 95%, which is 0.32 m (accounting for 118.5%) difference from the 521 LNWL estimated by the proposed framework. When the guaranteed rate is 98%, the LNWL 522 obtained using the NSDC method is 0.16 m at Shijiao station, which is close to the LNWL 523 designed by our framework. The differences of the values are resulted from the ways to 524 consider the guaranteed rate and the steps of designing. In practical hydrological design, it is 525 contradictory that there are different LNWLs in the same river section. Therefore, it is 526 necessary to identify the advantages and limitations of each method, and select the suitable 527 method for practical design. 528

The NSDC method is simple in calculation, but it does not consider whether the water level process of the current is more uniform/uneven than that of the past. In theory, this lack of consideration would make the LNWLs estimated by the NSDC method inaccurate. According to the design principles, the variance variabilities in low water level series (the 3-, 10- and 15-day minimum water level series) of Gaodao station are overlooked, since the
annual average water level series does not contain the variability of variance. Meanwhile, the
abrupt changes of water levels at different time scales are unreasonably regarded as the same
at both stations.

To compare the water level duration curves designed by the two methods, the violin 537 plots are used to display their distributions. For Gaodao station, the shapes of the distributions 538 539 in the past and current are very similar (Fig.11(c)) in NSDC method, indicating that the uniformity of water level process hardly changes over time. However, by using our 540 comprehensive framework, the water level distribution of the past environment has a longer 541 upper tail (Fig. 11 (d)), indicating that the past water level process has a number of high water 542 levels far from the average. In other words, the water level process in the past is more uneven 543 than that in the current. Therefore, only the duration curves designed by comprehensive 544 framework are match with the detection results of intra-annual variability. For Shijiao station, 545 because the distribution of the past has both long upper and lower tails, the synthetic duration 546 curve designed by NSDC method for the past environment is more uneven than that of the 547 current environment (Fig. 11 (e)). This indication is contrary to the detection results. However, 548 the duration curves designed by our comprehensive framework show that the water level 549 process in the current is more uneven (Fig. 11 (f)), which is match with the detection results. 550 To sum up, in NSDC method, the variabilities in the annual water level process are 551 552 generalized by the variability of annual average water level. When the annual average water level series has variance variability, it is easy to cause some water levels (especially high 553

554 water levels and low water levels) to be over enlarged or reduced when reconstructing the series, resulting in unreasonable duration curves for estimations. However, if the water level 555 variabilities are consistent across the timescales and the intra-annual variability is not 556 significant, the NSDC method is more convenient than the developed framework, and the 557 design results are similar. In our comprehensive framework, the variabilities are considered 558 comprehensively, therefore, the estimated LNWLs are more reliable when the waterway is 559 560 faced with complex multi-scale variabilities and prominent intra-annual variability. As for the future LNWLs, according to the expressions of deterministic components and the 561 formulas for "decomposition and composition", the framework can compose the "stationary" 562 series under any year  $t_0$  of the future. Then, the LNWLs can be estimated according to the 563 original steps. Since the forms of variabilities in the case study are all abrupt changes, the 564 LNWLs in the current environment can be taken as the LNWLs of the future. However, with 565 the extension of water level series, the deterministic components are variable. Therefore, the 566 LNWLs estimated in this way only have the significance when the external environment 567 remains relatively stable (e.g., the operating mode of the reservoir remains unchanged, the 568 riverbed has no obvious undercutting). 569

# 570 **6. Conclusion**

In this paper, we proposed a comprehensive framework for designing LNWLs by considering the intra- and inter-annual variabilities in water levels, as well as the shipping risks during the annual and multi-annual periods. Since the hydrological variability has extended to higher-order moments, the variability in variance is also described when 575 decomposing the observed series and is eliminated when composing the new series.

The proposed framework was illustrated by case studies of Gaodao and Shijiao 576 monitoring stations in the North River basin, South China. The results showed that water 577 levels at multiple timescales at Gaodao and Shijiao stations had different degrees of abrupt 578 changes in their means and variances. Meanwhile, the abrupt changes in  $C_d$ ,  $C_n$ ,  $C_c$  and GI579 indicated that the shape of annual water level process changed with time. By taking the inter-580 581 and intra-annual variability into consideration, the LNWLs of Gaodao station were recommended as 22.32 m (for 95% guaranteed rate) and 21.84 m (for 98% guaranteed rate), 582 and the LNWLs of Shijiao station were recommended as 0.27 m (for 95% guaranteed rate) 583 and 0.15 m (for 98% guaranteed rate). 584

To evaluate the impact of variance variability on the LNWL, the LNWLs estimated by the proposed framework without considering any variability in variance were used for comparison. Results indicated that the variability in variance had impact on the estimation of LNWLs, so this variability should not be ignored. Additionally, the magnitude of the impacts mainly depended on the extent of variance variability and the difference between the water level values of the two reconstructed series at assumed exceedance probability.

To understand the differences in the estimated LNWLs caused by different design methods, LNWLs designed by the NSDC method (adding the consideration of variance variability) were compared with the recommended LNWLs of the proposed framework. The differences in the LNWLs between the two methods at Gaodao station were -0.07 m (95% guaranteed rate) and 0.24 m (98% guaranteed rate); at Shijiao station, the differences in the 596 LNWLs were 0.32 m (95% guaranteed rate) and 0.01 m (98% guaranteed rate). In theory, 597 some daily water levels (especially high- and low-water levels) could be over enlarged or 598 reduced when reconstructing the pre- and post-impact series, resulting in unreasonable 599 duration curves for estimations. Evidence also showed that the reconstructed series as well 600 as the duration curves calculated by the NSDC method did not match the detected intra-601 annual variabilities.

In summary, our comprehensive framework is suitable for waterways faced with complicated water level variabilities, especially where the variabilities of water levels at different time scales are different or where the historical data contain various kinds of variations. In future research, this framework can be applied in more sites, and the uncertainty of estimated LNWLs caused by the selection of typical years can be further studied.

607 **Declarations** 

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612 **Conflicts of interest** 

613 Authors have no conflicts of interest to disclose.

# 614 Availability of data and material

The authors gratefully acknowledged the valuable hydrological data and informationprovided by the Hydrology Bureau of Guangdong Province, China. The data that support the

617 findings of this study are available from the corresponding author upon reasonable request.

#### 618 **Code availability**

- 619 The code is available from the corresponding author upon reasonable request.
- 620 Authors' contributions
- Ping Xie developed the main ideas. Lu Wang implemented the algorithms of the
- methods. Ping Xie collected the data used in the case study. Ping Xie, Chong-Yu Xu and Yan-
- Fang Sang provided funding for this study. Lu Wang wrote the original draft of the manuscript.
- 624 Chong-Yu Xu, Yan-Fang Sang and Jie Chen supervised this study, reviewed and edited the
- original draft. Tao Yu validated the results of this study.

#### 626 **Ethics approval**

627 Not applicable.

# 628 **Consent to participate**

629 Not applicable.

#### 630 **Consent for publication**

631 We confirm that this work is original and has not been published elsewhere, nor is it

632 currently under consideration for publication elsewhere. All authors have read and approved

the manuscript being submitted and agree to its publication in this journal.

# 634 **References**

- 635 Afifi AA, Azen SP (1972) Statistical analysis, a computer oriented approach. Academic
- 636 Press, Harcourt Brace Jovanonich Publishers, New York, 366.
- 637 Ahn KH, Palmer RN (2016) Use of a nonstationary copula to predict future bivariate low

- flow frequency in the Connecticut river basin. *Hydrological Processes*, 30(19), 3518-
- 639 3532. https://doi.org/10.1002/hyp.10876
- 640 Bartels R (1982) The rank version of von Neumann's ratio test for randomness. Journal of
- 641 *the American Statistical Association*, 77(377), 40-46. https://
   642 doi.org/10.1080/01621459.1982.10477764
- Bras RL, Rodriguez-Iturbe I (1993) *Random functions and hydrology*. Dover Publications,
  New York.
- Brown MB, Forsythe AB (1974) Robust tests for the equality of variances. *Journal of the American Statistical Association*, 69(346), 364-367.
- 647 Cai SY, Lei XH, Meng XY, Yi J, Mahalingam S, Gao XZ, Hamed VN (2018) Ecological
- 648 flow analysis method based on the comprehensive variation diagnosis of Gini
- 649 coefficient. Journal of Intelligent and Fuzzy Systems, 34(2), 1025-1031.
- 650 https://doi.org/10.3233/JIFS-169396
- 651 Cannon AJ (2010) A flexible nonlinear modelling framework for nonstationary generalized
- extreme value analysis in hydroclimatology. *Hydrological Processes*, 24(6),
  https://doi.org/10.1002/hyp.7506
- 654 Christodoulou A, Christidis P, Bisselink B (2020) Forecasting the impacts of climate change
- on inland waterways. *Transportation Research Part D: Transport and Environment*, 82,
- 656 102159. https://doi.org/10.1016/j.trd.2019.10.012
- 657 Cleveland RB, Cleveland WS, McRae JE, Terpenning I (1990) STL: A seasonal trend
- decomposition procedure based on loess. *Journal of Official Statistics*, 6(1), 3–73.

- Cui T, Tian FQ, Yang T, Wen J, Khan MYA (2020) Development of a comprehensive
  framework for assessing the impacts of climate change and dam construction on flow
  regimes. *Journal of Hydrology*, 590, 125358.
  https://doi.org/10.1016/j.jhydrol.2020.125358
- Dai SB, Yang SL, Cai AM (2008) Impacts of dams on the sediment flux of the Pearl River,
- southern China. *CATENA*, 76(1): 36-43. https://doi.org/10.1016/j.catena.2008.08.004
- 665 Du JK, Wu XS, Wang ZL, Li J, Chen XH (2020) Reservoir-Induced Hydrological Alterations
- 666 Using Ecologically Related Hydrologic Metrics: Case Study in the Beijiang River,
- 667 China. *Water*, 12(7). https://doi.org/10.3390/w12072008
- 668 Feng Y, Shi P, Qu S, Mou SY, Chen C, Dong FC (2020) Nonstationary flood coincidence
- risk analysis using time-varying copula functions. *Scientific Reports*, 10(1), 3395.

670 https://doi.org/10.1038/s41598-020-60264-3

- Gado TA, Nguyen VTV (2016a) An at-site flood estimation method in the context of
- nonstationarity I. A simulation study. *Journal of Hydrology*, 535, 710-721.
- 673 https://doi.org/10.1016/j.jhydrol.2015.12.063
- Gado TA, Nguyen VTV (2016b) An at-site flood estimation method in the context of
- 675 nonstationarity II. Statistical analysis of floods in Quebec. *Journal of Hydrology*, 535,
- 676 722–736. https://doi.org/10.1016/j.jhydrol.2015.12.064
- 677 Gau HS, Chen TC, Chen JS, Liu CW (2007) Time series decomposition of groundwater level
- 678 changes in wells due to the Chi-Chi earthquake in Taiwan: a possible hydrological
- 679 precursor to earthquakes. *Hydrological Processes*, 21, 510-524.

- 680 https://doi.org/10.1002/hyp.6257
- Gumbel EJ (1954) Applications of the circular normal distribution. Journal of the American 681 49(266), 267-297. 682 **Statistical** Association, https://doi.org/10.1080/01621459.1954.10483505 683 Guttman NB, Plantico MS (1989) On an additive model of daily temperature climates. 684 https://doi.org/10.1175/1520-Journal ofClimate, 2(10), 1207-1209. 685 686 0442(1989)002<1207:OAAMOD>2.0.CO;2
- 687 He W, Lian JJ, Zhang J, Yu XD, Chen S (2019) Impact of intra-annual runoff uniformity and
- global warming on the thermal regime of a large reservoir. *Science of Total Environment*,

689 658, 1085-1097. https://doi.org/10.1016/j.scitotenv.2018.12.207

690 Hu YM, Liang ZM, Singh VP, Zhang XB, Wang J, Li BQ, Wang HM (2018) Concept of

691 equivalent reliability for estimating the design flood under non-stationary conditions.

- 692 Water Resources Management, 32, 997-1011. https://doi.org/10.1007/s11269-017-
- 693 1851-у
- Hu YM, Liang ZM, Jiang XL, Bu H (2015) Non-stationary hydrological frequency analysis
- based on the reconstruction of extreme hydrological series. Proceedings of the
- 696 International Association of Hydrological Sciences 371, 163-166. https://doi.org/
- 697 10.5194/piahs-371-163-2015
- <sup>698</sup> Jawitz JW, Mitchell J (2011) Temporal inequality in catchment discharge and solute export.
- 699 *Water Resources Research*, 47(10). https://doi.org/10.1029/2010WR010197.
- Jiang C, Xiong L, Xu CY, Guo SL (2015) Bivariate frequency analysis of nonstationary low-

- flow series based on the time-varying copula. *Hydrological Processes*, 29, 1521–1534.
  https://doi.org/10.1002/hyp.10288.
- Jiang C, Xiong LH, Yan L, Dong QJ, Xu CY (2019) Multivariate hydrologic design methods
- <sup>704</sup> under nonstationary conditions and application to engineering practice. *Hydrology and*
- 705 *Earth System Sciences*, 23, 1683–1704. https://doi.org/10.5194/hess-23-1683-2019
- Jonkeren O, Jourquin B, Rietveld P (2011) Modal-split effects of climate change: The effect
- of low water levels on the competitive position of inland waterway transport in the river
- Rhine area. Transportation Research Part A Policy & Practice, 45(10), 1007-1019.
- 709 https://doi.org/10.1016/j.tra.2009.01.004
- Jonkeren O, Rietveld P, Ommeren JV, Linde AT (2014) Climate change and economic
- consequences for inland waterway transport in Europe. *Regional Environmental*

712 *Change*, 14, 953–965. https://doi.org/10.1007/s10113-013-0441-7

- 713 Kendall MG (1955) *Rank correlation methods*. Griffin, London.
- Kling GW, Hayhoe K, Johnson LB, Magnuson JJ, Polassky S, Robinson SK, Shuter BJ, et
- al. (2003) Confronting climate change in the great lakes region: Impacts on our
- *communities and ecosystems*. https://www.researchgate.net/publication/248822899
- Lee AFS, Heghinian SM (1977) A shift of the mean level in a sequence of independent
- normal random variable: A Bayesian approach. *Technometrics*, 19(4), 503-506.
- 719 https://doi.org/10.1080/00401706.1977.10489592
- Li GF, Xiang XY, Guo CX (2016) Analysis of nonstationary change of annual maximum
- level records in the Yangtze river estuary. *Advances in Meteorology*, 2016, 1-14.

722 http://dx.doi.org/10.1155/2016/7205723

- Li M, Zhang T, Feng P (2019) A nonstationary runoff frequency analysis for future climate
- change and its uncertainties. *Hydrological Processes*, 33(21), 2759-2771.
  https://doi.org/10.1002/hyp.13526
- Li R, Tang CY, Li X, Jiang T, Shi YP, Cao YJ (2019) Reconstructing the historical pollution
- <sup>727</sup> levels and ecological risks over the past sixty years in sediments of the Beijiang River,
- South China. Science of The Total Environment, 649: 448-460.
   https://doi.org/10.1016/j.scitotenv.2018.08.283
- Liang ZM, Yang J, Hu YM, Wang J, Li BQ, Zhao JF (2017) A sample reconstruction method
- based on a modified reservoir index for flood frequency analysis of non-stationary
- hydrological series. Stochastic Environmental Research and Risk Assessment, 32, 1561-
- 733 1571. https://doi.org/10.1007/s00477-017-1465-1
- 734 Linde F, Ouahsine A, Huybrechts N, Sergent P (2017) Three-dimensional numerical
- simulation of ship resistance in restricted waterways: Effect of ship sinkage and channel
- restriction. Journal of Waterway, Port, Coastal and Ocean Engineering, 143(1), 1-11.
- 737 https://doi.org/10.1061/(ASCE)WW.1943-5460.0000353
- Liu Y, Hao YH, Fan YH, Wang TK, Liu YC, Jim-Yeh TC (2014) A nonstationary extreme
- value distribution for analysing the cessation of karst spring discharge. *Hydrological*
- 740 *Processes*, 28(20), https://doi.org/10.1002/hyp.10013
- 741 López-Moreno JI, Vicente-Serrano SM, Zabalza J, Beguer'ia S, Lorenzo-Lacruz J, Azorin-
- 742 Molina C, Morán-Tejeda E (2013) Hydrological response to climate variability at

- different time scales: A study in the Ebro basin. *Journal of Hydrology*, 477,
- 744 https://doi.org/175-188. 10.1016/j.jhydrol.2012.11.028
- 745 Lu XX, Zhang SR, Xie SP, Ma PK (2007) Rapid channel incision of the lower Pearl River
- 746 (China) since the 1990s as a consequence of sediment depletion. *Hydrology and Earth*
- 747 System Sciences, 11: 1897-1906. https://doi.org/10.5194/hess-11-1897-2007
- Luo Y, Liu S, Fu SL, Liu JS, Wang GQ, Zhou GY (2008) Trends of precipitation in Beijiang
- River Basin, Guangdong Province, China. *Hydrological Processes*, 22, 2377-2386.
- 750 https://doi.org/10.1002/hyp.6801
- Machiwal D, Jha MK (2012) *Hydrologic time series analysis: Theory and practice*, Springer
  Netherlands.
- Magilligan FJ, Graber BE (1996) Hydroclimatological and geomorphic controls on the
   timing and spatial variability of floods in New England, USA. *Journal of Hydrology*,

755 178, 159-180. https://doi.org/10.1016/0022-1694(95)02807-2

- 756 Magilligan FJ, Nislow KH (2005). Changes in hydrologic regime by dams. *Geomorphology*,
- 757 71(1), 61-78. https://doi.org/10.1016/j.geomorph.2004.08.017
- 758 Maidment DR (1993) Handbook of hydrology, McGraw-Hill. New York.
- Mann HB (1945) Nonparametric tests against trend. *Econometrica*, 13, 245-259.
  https://doi.org/10.2307/1907187
- 761 Masaki Y, Hanasaki N, Takahashi K, Hijioka Y (2014) Global-scale analysis on future
- changes in flow regimes using Gini and Lorenz asymmetry coefficients. *Water*
- 763 *Resources Research*, 50(5), 4054-4078. https://doi.org/10.1002/2013WR014266

- 764 Milly PCD., Betancourt J, Falkenmark M., Hirsch RM, Zbigniew W, Lettenmaier DP,
- 765 Stouffer RJ (2008) Stationarity is dead: Whither water management? *Science* 319, 573-
- 766 574. https://doi.org/10.1126/science.1151915
- 767 Milly PCD, Betancourt J, Falkenmark M, Hirsch RM, Kundzewicz ZW, Lettenmaier DP,
- Stouffer RJ, et al. (2015) On critiques of "stationarity is dead: Whither water
  management?". *Water Resources Research*, 51(9), 7785-7789.
  https://doi.org/10.1002/2015WR017408
- 771 Ministry of Development of the People's Republic of China, General Administration of
- 772 Quality Supervision & Inspection and Quarantine of the People's Republic of China
- (2014) *Navigation standard in inland river (GB50139-2014)*, China Plan Press, Beijing.
- 774 (in Chinese)
- Montanari A, Koutsoyiannis D (2014) Modeling and Mitigating Natural Hazards:
  Stationarity is Immortal! *Water Resources Research*, 50(12).
  https://doi.org/10.1002/2014WR016092
- Oztanriseven F, Nachtmann H (2017). Economic impact analysis of inland waterway
  disruption response. *The Engineering Economist*, 62(1), 73-89.
  https://doi.org/10.1080/0013791X.2016.1163627
- 781 Palleiro L, Rodríguez-Blanco ML, Taboada-Castro MM (2014) Hydrological response of a
- humid agroforestry catchment at different time scales. *Hydrological Processes*, 28(4),
- 783 1677-1688. https://doi.org/10.1002/hyp.9714
- Ren K., Huang S, Huang Q, Wang H, Leng GY (2018) Environmental Flow Assessment

- 785 Considering Inter- and Intra-Annual Streamflow Variability under the Context of Non-
- 786 Stationarity. *Water*, 10, 1737. https://doi.org/10.3390/w10121737
- 787 Rootzen H, Katz RW (2013) Design Life Level: Quantifying risk in a changing climate.
- 788 Water Resources Research, 49, 5964-5972. https://doi.org/10.1002/wrcr.20425
- 789 Rosner A, Vogel RM, Kirshen PH (2014) A risk based approach to flood management
- decisions in a nonstationary world. *Water Resources Research*, 50(3), 1928-1942.
- 791 https://doi.org/10.1002/2013WR014561
- 792 Salas JD, Obeysekera J (2014) Revisiting the concepts of return period and risk for
- nonstationary hydrologic extreme events. *Journal of Hydrologic Engineering*, 19, 554-

794 568. https://doi.org/10.1061/(ASCE)HE.1943-5584.0000820

- 795 Sarhadi A, Burn DH, Ausín MC, Wiper MP (2016) Time-varying nonstationary multivariate
- risk analysis using a dynamic Bayesian copula, *Water Resources Research*, 52, 2327–
- 797 2349. https://doi.org/10.1002/2015WR018525
- 798 Sheskin D (2011) Handbook of Parametric and Nonparametric Statistical Procedures, CRC
- 799 Press, USA.
- 800 Shi WL, Yu XZ, Liao WG, Wang Y, Jia BZ (2013) Spatial and temporal variability of daily
- 801 precipitation concentration in the Lancang River basin, China. *Journal of Hydrology*,
- 495, 197-207. https://doi.org/10.1016/j.jhydrol.2013.05.002
- 803 Shiau JT, Wu FC (2007) Pareto optimal solutions for environmental flow schemes
- incorporating the intra annual and interannual variability of the natural flow regime.
- 805 *Water Resources Research*, 43 (6). https://doi.org/10.1029/2006WR005523

- Singh KP, Sinclair RA (1972) Two distribution method for flood frequency analysis. *Journal of Hydraulics Division*, 98(1), 29-44.
- 808 Singh VP, Wang SX, Zhang L (2005) Frequency analysis of nonidentically distributed
- hydrologic flood data. *Journal of Hydrology*, 307(1–4), 175-195.
  https://doi.org/10.1016/j.jhydrol.2004.10.029
- 811 Statsoft Inc. (2003) *Statistica: The Small Book. Statsoft Inc.*, Tulsa, 144.
- 812 Stojković M, Kostić S, Plavšić J, Prohaska S (2017) A joint stochastic-deterministic approach
- for long-term and short-term modelling of monthly flow rates. *Journal of Hydrology*,
- 814 544.555S. https://doi.org/10.1016/j.jhydrol.2016.11.025
- 815 Strupczewski WG, Singh VP, Feluch W (2001a) Non-stationary approach to at-site flood
- frequency modelling I. Maximum likelihood estimation. *Journal of Hydrology*, 248,
- 817 123-142. https://doi.org/10.1016/S0022-1694(01)00397-3
- 818 Strupczewski WG, Kaczmarek Z (2001) Non-stationary approach to at-site flood frequency
- modelling II. Weighted least squares estimation. *Journal of Hydrology*, 248, 143-151.
- 820 https://doi.org/10.1016/S0022-1694(01)00398-5
- 821 Strupczewski WG, Singh VP, Mitosek HT (2001b) Non-stationary approach to at-site flood
- frequency modelling III. Flood analysis of Polish rivers. *Journal of Hydrology*, 248,
- 823 152-167. https://doi.org/10.1016/S0022-1694(01)00399-7
- Tu X, Singh VP, Chen XH, Chen L, Zhang Q, Zhao Y (2015) Intra-annual distribution of
- streamflow and individual impacts of climate change and human activities in the
- Dongijang river basin, China. Water Resources Management, 29(8), 2677-2695.

#### https://doi.org/10.1007/s11269-015-0963-5

Valle D, Kaplan D (2019). Quantifying the impacts of dams on riverine hydrology under 828 non-stationary conditions using incomplete data and Gaussian copula models. Science 829 The Total 830 of Environment, 677. 599-611. https://doi.org/10.1016/j.scitotenv.2019.04.377 831 Villarini G, Smith JA, Serinaldi F, Bales JD, Bates PD and Krajewski WF (2009) Flood 832 833 frequency analysis for nonstationary annual peak records in an urban drainage basin. Advances in Water 834 Resources. 32(8):1255-1266. https://doi.org/10.1016/j.advwatres.2009.05.003 835 Vinnikov KY, Robock A (2002) Trends in moments of climatic indices. Geophysical 836 Research Letters, 29(2), 1027. https://doi.org/10.1029/2001GL014025 837 Vogel RM, Yaindl C, Walter M (2011) Nonstationarity: Flood Magnification and Recurrence 838 Reduction Factors in the United States. Journal of the American Water Resources 839 Association, 47(3): 464-474. https://doi.org/10.1111/j.1752-1688.2011.00541.x 840 Wang D, Ding H, Singh VP, Shang XS, Liu DF, Wang YK, Zeng XK, et al. (2015) A hybrid 841 wavelet analysis-cloud model data-extending approach for meteorologic and 842 hydrologic time series. Journal of Geophysical Research Atmospheres, 120, 4057–4071. 843 https://doi.org/10.1002/2015JD023192 844 Wang Y, Chen X, Borthwick AGL, Li TH, Liu HH, Yang SF, Zheng CM, et al. (2020) 845 846 Sustainability of global golden inland waterways. Nature Communications, 11, 1553. https://doi.org/10.1038/s41467-020-15354-1 847

848	Waylen P, Woo MK	(1982) Predictio	on of annual floods	generated b	y mixed	processes.	Water
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- 849 *Resources Research*, 18 (4), 1283-1286. https://doi.org/10.1029/WR018i004p01283
- 850 Willems JJ, Busscher T, Woltjer J, Arts J (2018) Co-creating value through renewing
- 851 waterway networks: a transaction-cost perspective. *Journal of Transport Geography*,
- 69, 26–35. https://doi.org/10.1016/j.jtrangeo.2018.04.011
- 853 Wu CH, Huang GR, Yu HJ, Chen ZJ, Ma JG (2014) Impact of Climate Change on Reservoir
- Flood Control in the Upstream Area of the Beijiang River Basin, South China. *Journal*
- *of Hydrometeorology*, 15(6), 2203-2218. https://doi.org/10.1175/JHM-D-13-0181.1
- Xie P, Wu ZY, Sang YF, Gu HT, Zhao YX, Singh VP (2018) Evaluation of the significance
- of abrupt changes in precipitation and runoff process in China. *Journal of Hydrology*,
- 858 560, 451-460. https://doi.org/10.1016/j.jhydrol.2018.02.036
- 859 Yan L, Xiong LH, Liu DD, Hu TS, Xu CY (2016) Frequency analysis of nonstationary annual
- 860 maximum flood series using the time varying two component mixture distributions.
- 861 *Hydrological Processes*, 31(1), 69-89. https://doi.org/10.1002/hyp.10965
- 862 Yang YP, Zhang MJ, Liu WL, Wang JJ, Li XX (2019) Relationship between Waterway
- 863 Depth and Low-Flow Water Levels in Reaches below the Three Gorges Dam. *Journal*
- 864 of Waterway, Port, Coastal, and Ocean Engineering, 145(1), 04018032.
- 865 https://doi.org/10.1061/(ASCE)WW.1943-5460.0000482
- 866 Yao LL, Libera DA, Kheimi M, Sankarasubramanian A, Wang DB (2020) The Roles of
- 867 Climate Forcing and its Variability on Streamflow at Daily, Monthly, Annual, and
- 868 Long term Scales. *Water Resources Research*.

https://doi.org/10.1029/2020WR027111

- Yevjevich V (1972) *Stochastic processes in hydrology*. Water Resources Publications, Fort
  Collins, Colorado, USA.
- 872 Zhao JY, Xie P, Zhang MY, Sang YF, Chen J, Wu ZY (2018) Nonstationary statistical
- approach for designing LNWLs in inland waterways: A case study in the downstream
- of the Lancang River. Stochastic Environmental Research and Risk Assessment, 32,
- 875 3273-3286. https://doi.org/10.1007/s00477-018-1606-1
- 876 Zheng F, Tao R, Maier HR, See LM, Savic D, Zhang TQ, Chen QW, et al. (2018)
- 877 Crowdsourcing methods for data collection in geophysics: State of the art, issues, and
- future directions, *Reviews of Geophysics*, 56, https://doi.org/10.1029/2018RG000616
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# 894 Six Tables submitted

# Table 1 Detection results of inter-annual variability

Gaodao station		3-day	10-day	15-day	30-day	60-day	90-day
Maan	Form of variability	Abrupt	Abrupt	Abrupt	Abrupt	Abrupt	Abrupt
Ivieali	Form of variability	change	change	change	change	change	change
Mean	Position of change points	1999↑	1999↑	1999↑	1998↑	1998↑	1998↑
Variance	Form of variability	Abrupt	Abrupt	Abrupt			
v allallee	Point of Variability	change	change	change	-	-	-
Variance	Position of change points	1999↑	2002↑	2002↑	-	-	-
	Gaodao station	120-dav	150-dav	180-dav	240-dav	300-dav	Annual
	Gaodao station	120-uay	150-uay	180 <b>-</b> uay	240-uay	500-uay	average
Mean	Form of variability	Abrupt	Abrupt	Abrupt	Abrupt	Abrupt	Abrupt
Ivicali	Form of variability	change	change	change	change	change	change
Mean	Position of change points	1998↑	1998↑	1998↑	1998↑	1998↑	1998↑
Variance	Form of variability	-	-	-	-	-	-
Variance	Position of change points	-	-	-	-	-	-
Shijiao station		3-day	10-day	15-day	30-day	60-day	90-day
Moon	Form of variability	Abrupt	Abrupt	Abrupt	Abrupt	Abrupt	Abrupt
Ivicali		change	change	change	change	change	change
Mean	Position of change points	2006↓	2006↓	2006↓	2006↓	2005↓	2005↓
Variance	Form of variability	Abrupt	Abrupt	Abrupt	Abrupt	Abrupt	Abrupt
variance		change	change	change	change	change	change
Variance	Desition of change points	2003↑;	2003†;	2003↑;	2003†;	2003↑;	2003↑;
variance	I osition of change points	2007↓	2007↓	2007↓	2007↓	2007↓	2007↓
	Shijiao station	120 day	150 day	180 day	240 day	300 day	Annual
	Shijiao station	120-uay	150-uay	180 <b>-</b> uay	240-uay	500-uay	average
Maan	Form of variability	Abrupt	Abrupt	Abrupt	Abrupt	Abrupt	Abrupt
Ivicali	1 on in variability	change	change	change	change	change	change
Mean	Position of change points	2005↓	2005↓	2005↓	2005↓	2007↓	2007↓
Varianco	Form of variability	Abrupt	Abrupt	Abrupt	Abrupt	Abrupt	Abrupt
v arrance	ronn or variability	change	change	change	change	change	change
Variance	Position of change points	2003†;	2003†;	2003†;	2003↑	2003†;	2003†;

Note: The test is taken under the significance level of  $\alpha = 0.05$ ; "-" means there is no significant variability; " $\uparrow$ " or " $\downarrow$ "

means that the abrupt change is upward or downward.

#### 

# Table 2 Detection results of intra-annual variability

Station	ion Indicator		$C_n$	$C_c$	GI
	E	Abrupt	Abrupt	Abrupt	Abrupt
Gaodao	Form of variability	change	change	change	change
	Position of change points	1998↓	1998↓	1998↓	1998↑
	Form of variability	Abrupt	Abrupt	Abrupt	Abrupt
Shijiao	Form of variability	change	change	change	change
	Position of change points	2006↑	2007↑	2007↑	2007↓

the abrupt change is downward.

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Table 3 Estimated parameters for pre- and post-impact periods

Gaod	Gaodao station		10-day	15-day	30-day	60-day	90-day
EV/	Pre-impact	19.98	20.04	20.05	20.08	20.15	20.22
EX/m	Post-impact	22.23	22.55	22.59	22.88	23.19	23.33
C	Pre-impact	0.011	0.011	0.011	0.013	0.011	0.012
$C_{v}$	Post-impact	0.037	0.040	0.040	0.013	0.011	0.011
C	Pre-impact	0.392	0.065	0.265	0.652	0.620	0.519
$\mathcal{L}_s$	Post-impact	0.392	0.065	0.265	0.652	0.620	0.519
Cood	a a station	120 day	150 day	100 day	240 day	200 day	Annual
Gaou	ao station	120-day	150-day	180-day	240-day	500-day	average
EV/m	Pre-impact	20.28	20.34	20.42	20.65	20.96	21.20
$L\Lambda/111$	Post-impact	23.38	23.43	23.47	23.53	23.59	23.63
C	Pre-impact	0.012	0.012	0.012	0.013	0.014	0.013
$C_{v}$	Post-impact	0.011	0.011	0.011	0.012	0.012	0.012
C	Pre-impact	0.625	0.454	0.317	0.097	0.401	0.364
$C_s$	Post-impact	0.625	0.454	0.317	0.097	0.401	0.364
Shijia	Shijiao station		10-day	15-day	30-day	60-day	90-day
EV/m	Pre-impact	4.48	4.55	4.58	4.65	4.75	4.84
	Post-impact	0.66	0.76	0.832	0.89	1.16	1.22
C	Pre-impact	0.058	0.059	0.060	0.065	0.059	0.064
$C_{v}$	Post-impact	0.941	0.803	0.718	0.723	0.613	0.596
C	Pre-impact	0.164	0.229	0.389	0.385	0.708	0.772
$C_s$	Post-impact	0.184	0.246	0.389	0.385	0.708	0.772
Shiji	a station	120 day	150 day	180 day	$240  \mathrm{day}$	200 day	Annual
Sinjiao station		120-uay	150-uay	180-uay	240-uay	500-uay	average
FV/m	Pre-impact	4.93	5.02	5.12	5.44	5.77	6.03
	Post-impact	1.32	1.40	1.49	1.79	2.15	2.42
C	Pre-impact	0.066	0.068	0.068	0.066	0.066	0.062
υv	Post-impact	0.556	0.540	0.534	0.638	0.442	0.382
C	Pre-impact	0.638	0.658	0.538	0.546	0.820	0.770
$C_S$	Post-impact	0.638	0.658	0.538	0.546	0.820	0.770

919	Note: the statistica	al parameter $EX$ represents the mean value, $C_V$ represents the mean value, $C_V$ represents the mean value of $C_V$ r	esents the n	on-uniformity	coefficient,	and $C_S$ represents
920	the coefficient of	skewness.				
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924		Table 4 Standards of desi	igning L	NWLs		
		Grade of channel	I~II	III~IV	V~VII	
		Annual guaranteed rate/%	99~98	98~95	95~90	
		Return period/year	10~5	5~4	4~2	
		Allowable carrying weight of a vessel/tons	2000	1000~500	300~50	
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Table 5 Designed water levels (in average) and amplifiers for scaling benchmark hydrographs

Water levels		Pre-impact period			Post-impact period		
		$\overline{Z_{d,j}}$ /m	$\overline{Z_{t,j}}$ /m	Amplifier	$\overline{Z_{d,j}}$ /m	$\overline{Z_{t,j}}$ /m	Amplifier
	Annual average	20.96	21.01	0.992	23.38	23.41	0.992
	3-day min	19.79	19.84	0.997	21.52	21.79	0.987
	10-day min	19.85	19.93	0.996	21.79	21.89	0.999
	15-day min	19.86	19.96	0.992	21.82	22.03	0.981
Gaodao	30-day min	19.87	19.96	0.996	22.63	22.61	1.010
station	60-day min	19.96	20.01	0.999	22.97	22.89	1.006
station	90-day min	20.01	20.04	1.001	23.11	23.13	0.990
	120-day min	20.07	20.12	0.995	23.16	23.17	1.002
	150-day min	20.13	20.18	0.996	23.20	23.25	0.991
	180-day min	20.21	20.28	0.994	23.24	23.27	1.003
	240-day min	20.43	20.43	1.010	23.30	23.28	1.007
	300-day min	20.71	20.73	0.994	23.34	23.34	0.997
	Annual average	5.71	5.73	1.027	1.63	1.38	1.161
	3-day min	4.26	4.25	1.004	0.13	0.05	2.822
	10-day min	4.33	4.32	1.000	0.24	0.20	1.052
	15-day min	4.35	4.35	0.998	0.32	0.26	1.305
Shijino	30-day min	4.39	4.37	1.008	0.34	0.26	1.376
station	60-day min	4.51	4.46	1.018	0.55	0.26	2.915
station	90-day min	4.58	4.52	1.014	0.60	0.32	1.567
	120-day min	4.65	4.60	1.011	0.69	0.50	0.924
	150-day min	4.73	4.69	0.997	0.75	0.52	1.661
	180-day min	4.82	4.71	1.094	0.81	0.54	1.778
	240-day min	5.13	5.04	1.005	0.81	0.71	0.669
	300-day min	5.45	5.52	0.902	1.33	1.12	1.239

# Table 6 Designed LNWLs under different considerations

Station	Considered	Annual guaranteed	Designed LNWLs/m		
Station	variability	rate /%	Past	Current	
	Maan and variance	95	19.84	22.32	
	Mean and variance	98	19.79	21.84	
	Moon only	95	19.87	22.43	
Gaodao	Weall only	98	19.81	22.32	
	Difference between	95	0.03	0.11	
	the two considerations	98	0.02	0.48	
	Maan and variance	95	4.41	0.27	
	Mean and variance	98	4.36	0.15	
Chilles	Maan anly	95	4.24	0.31	
Shijiao	Mean only	98	4.17	0.18	
	Difference between	95	-0.17	0.04	
	the two considerations	98	-0.19	0.03	

# Table 7 Designed LNWLs by NSDC method

Mathada		Considered veriabilities	Guaranteed	Designed LNWLs/m	
	wiethous	Considered variabilities	rate /%	Past	Current
	NSDC	Inter-annual variability (mean and	95	19.99	22.25
		variance)	98	19.83	22.08
Carlas	Proposed	Inter-annual variability (mean and	95	19.84	22.32
Gaodao	framework	variance) and intra-annual variability	98	19.79	21.84
	Differences between two methods/m		95	0.15	-0.07
			98	0.04	0.24
	NSDC	Inter-annual variability (mean and	95	3.04	0.59
		variance)	98	1.63	0.16
Chillon	Proposed	Inter-annual variability (mean and	95	4.41	0.27
Shijiao	framework	variance) and intra-annual variability	98	4.36	0.15
	Differences between two methods/m		95	-1.37	0.32
			98	-2.73	0.01

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# 981 Figure captions

- 982 Fig. 1 Framework for designing the LNWL
- 983 Fig. 2 The schematic diagram of Lorentz curve, section A and section B
- 984 Fig. 3 Different time scales of intra-annual water level process
- **Fig. 4** The North River basin in Guangdong Province, China
- 986 Fig. 5 (a) Variabilities in annual 3-, 15-, 120-day minimum water level series and annual
- 987 average water level series at Gaodao station (mean); (b) Variabilities in annual 3-, 15-,
- 988 120-day minimum water level series and annual average water level series at Shijiao989 station (mean).
- 990 Fig. 6 (a) Variabilities in annual 3-, 15-, 120-day minimum water level series and annual
- average water level series at Gaodao station (variance); (b) Variabilities in annual 3-,
- 992 15-, 120-day minimum water level series and annual average water level series at Shijiao
- station (variance).
- Fig. 7 (a) Monthly average water levels at Gaodao station; (b) Monthly average water levels
  at Shijiao station.
- 996 Fig. 8 (a) Designed intra-annual water level processes at Gaodao station; (b) Designed intra-

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annual water level processes at Shijiao station; (c) Water level duration curves at Gaodao station; (d) Water level duration curves at Shijiao station.

Fig. 9 Reconstructed annual 3-, 10- and 15-day minimum water level series: (a) Under the 999 past environment at Gaodao station; (b) Under the current environment at Gaodao 1000 station; (c) Under the past environment at Shijiao station; (d) Under the current 1001 environment at Shijiao station. 1002

- 1003 Fig. 10 Comparisons of low water levels estimated under different considerations: (a) Under the past environment at Gaodao station; (b) Under the current environment at Gaodao 1004 station; (c) Under the past environment at Shijiao station; (d) Under the current 1005 1006
- environment at Shijiao station.
- Fig. 11 (a) Duration curves at Gaodao station (by NSDC method); (b) Duration curves at 1007
- Shijiao station (by NSDC method); (c) Violin plots of duration curves at Gaodao station 1008
- (by NSDC method); (d) Violin plots of duration curves at Gaodao station (by proposed 1009
- framework); (e) Violin plots of duration curves at Shijiao station (by NSDC method); 1010
- (f) Violin plots of duration curves at Shijiao station (by proposed framework). 1011